

Modelling the Thermal Impact of a Repository for High-Level Radioactive Waste in a Clay Host Formation

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Thermal impact of the disposal of radioactive waste in clay

- **Geological disposal & problem specification**
 - General Context
 - Typical repository layout
 - The thermal issues associated with the disposal of heat-emitting wastes
- **T: Thermal evolution of a typical repository** ← **Simple, thermal model**
 - Typical temperature evolution
 - Model equation, implementation, results
- **T-H: Effect of / on groundwater flow** ← **Multiphysics model**
 - Thermo-hydraulic modelling of the far field
 - Model equations, implementation, results
- **Basic T-H-M: Uplift** ← **Multiphysics model**
 - Thermo-hydro-mechanical modelling of the far field
 - Model equations, implementation, results
- **Conclusions**

Geological disposal of long-lived, highly radioactive wastes

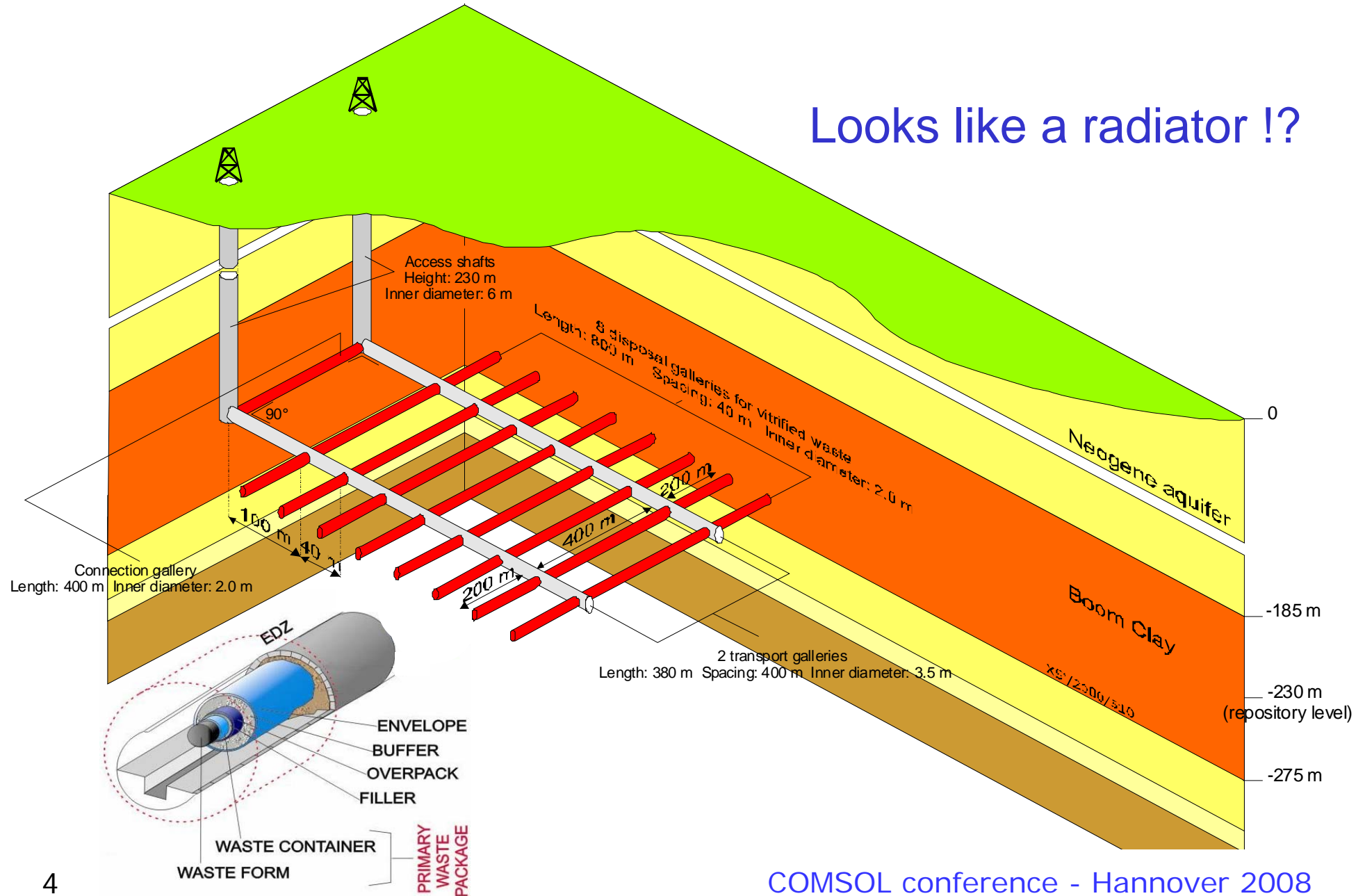
- **What can we do with our radioactive waste ?**
 - From nuclear power plants, medical, industrial activities
 - Main challenge = protection of men/environment during a **very long period of time** (10^4 ... 10^5 ... 10^6 years...)

- **Geological Disposal of high-level waste**
 - Accepted in a wide range of countries and by the EC
 - Engineered barriers + **geological barrier** : compatible with time scales associated with long-lived radioactive wastes:
 - Vitrified high-level waste (VHLW, reprocessed, COGEMA)
 - Spent fuel

- **Clays as potential hosts for a repository**
 - Very **low permeability** → solute transport by molecular diffusion
 - **Sorption** → delay and spread releases of radionuclides in time
 - If plastic clay: **self-sealing**, self-healing
 - Not a resource

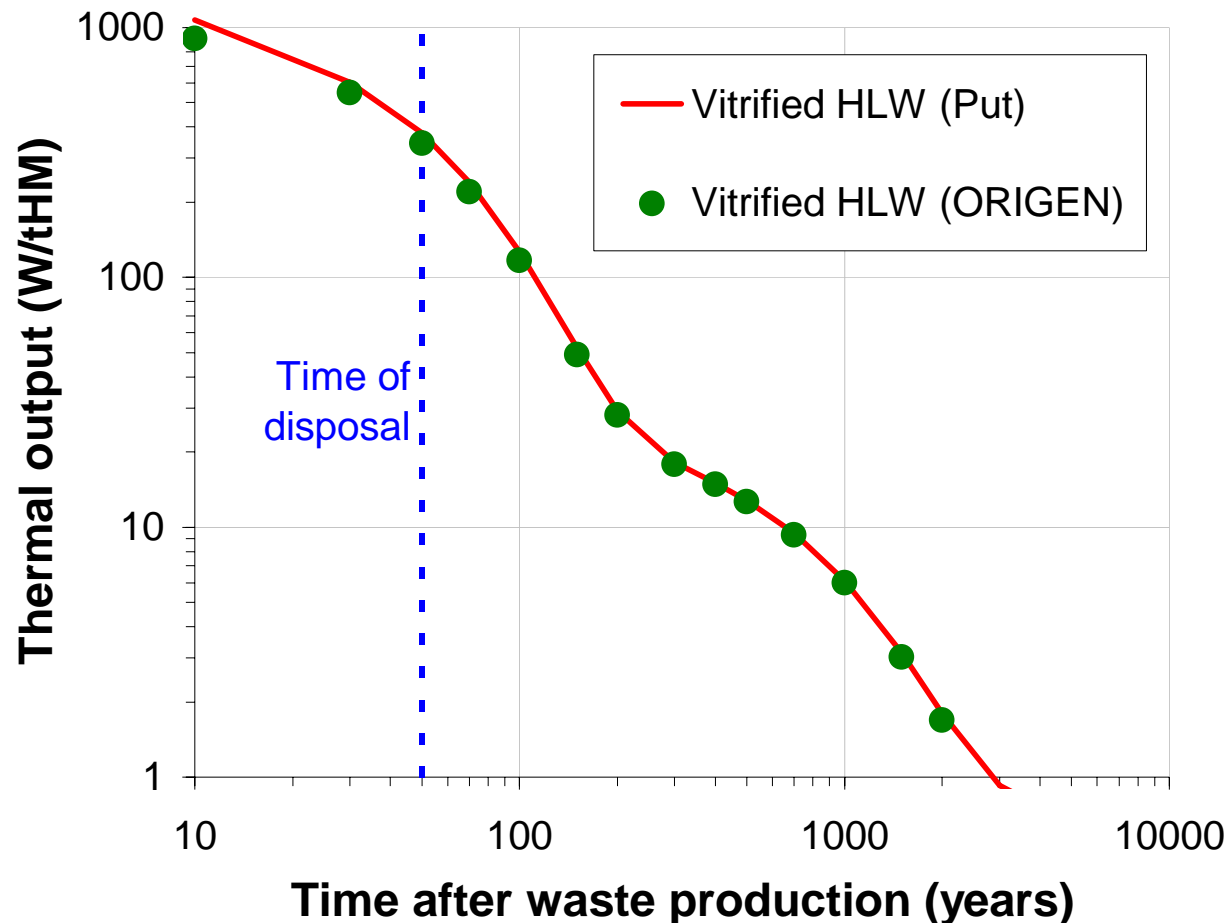
Typical repository design

Looks like a radiator !?



Some radioactive wastes can release considerable amounts of heat !

- Some radioactive wastes generate **a considerable amount of heat** due to radioactive decay, **even after interim storage (50-80 years)**
- Example: vitrified high-level waste (COGEMA)



Thermal issues associated with the disposal of heat-emitting waste

- How **hot** will it be ? ←
- Depends on waste type (radionuclide inventory)
- Engineered barriers & rock thermal properties
- Repository **design parameters** ←
- Disposal **galleries spacing**
- Waste **package pitch** within disposal galleries
- What could be the **consequences** of ΔT ?
- Chemical/geochemical ?
 - Thermal degradation of engineered barriers & waste forms ?
 - Solubility & migration parameters of radionuclides,... ?
 - Thermal decomposition of organic matter in Boom Clay, CO₂ ?
- Hydrogeology ?
 - **Far field: thermal impact on the aquifer** ? ←
- Mechanical ?
 - Near field: Thermo-Hydro-Mechanics of EBS, host rock ?
 - **Far field: uplift** ? ←

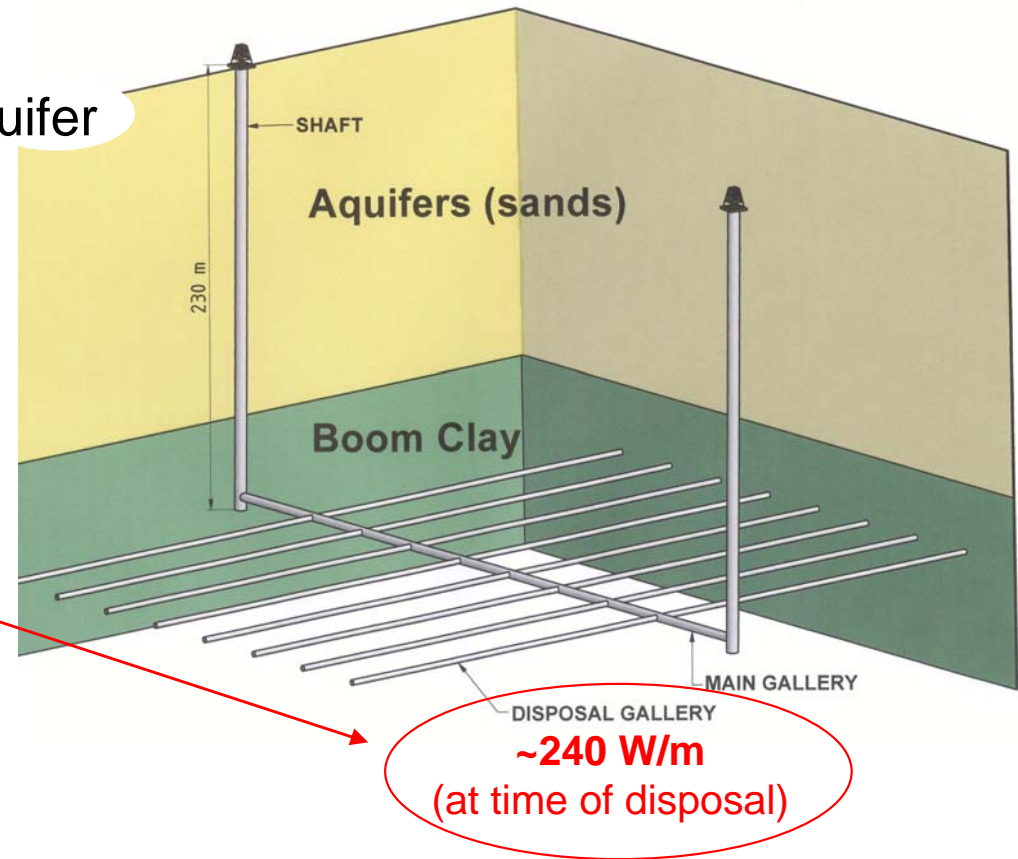
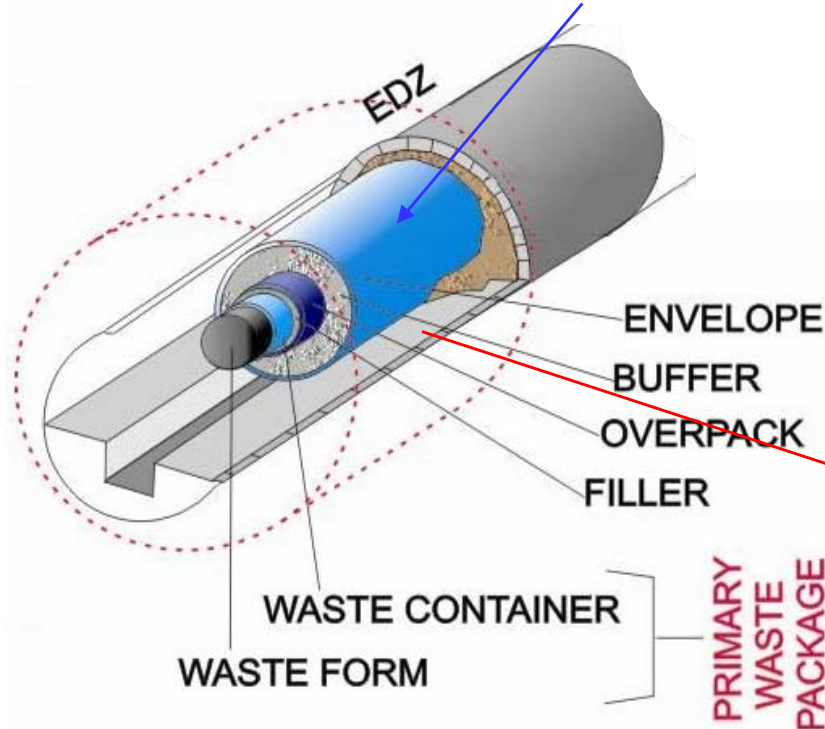
T: Thermal evolution of a typical repository (Belgian repository design)

Typical thermal loading for a disposal system:

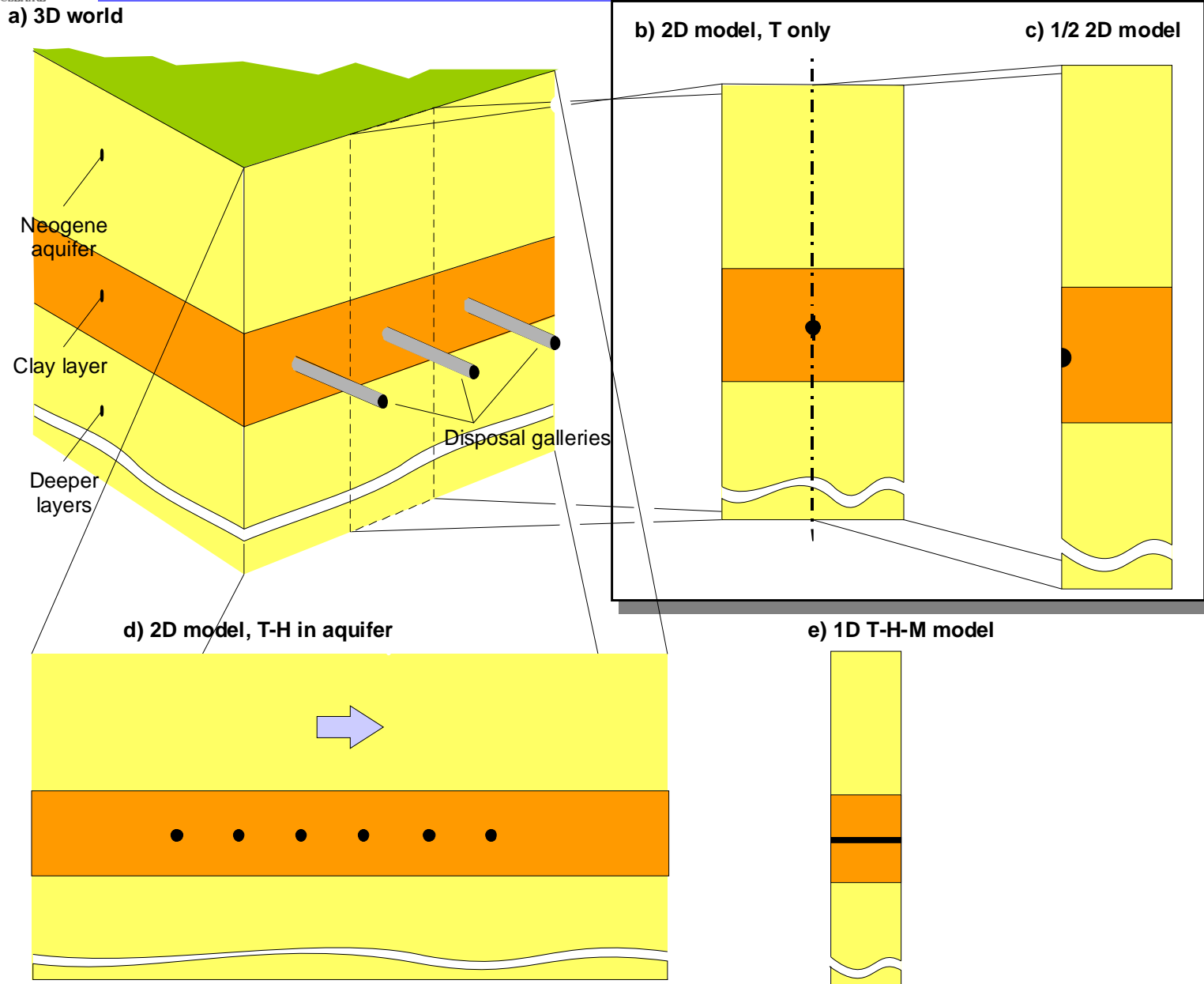
- VHLW: ~ **1 kW** per supercontainer after 60 years interim storage
- Supercontainer length = **4.2 m** (= package pitch: no spacing)
- Gallery spacing = **50 m**

Peak temperatures ?

- Conservative : no flow in aquifer



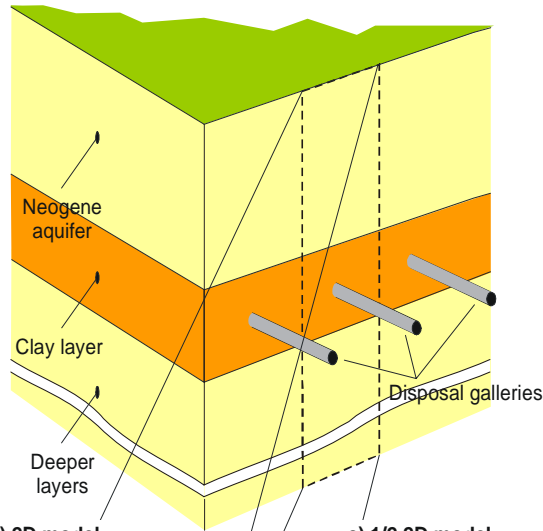
Reference geometry T, T-H & T-H-M model reduction



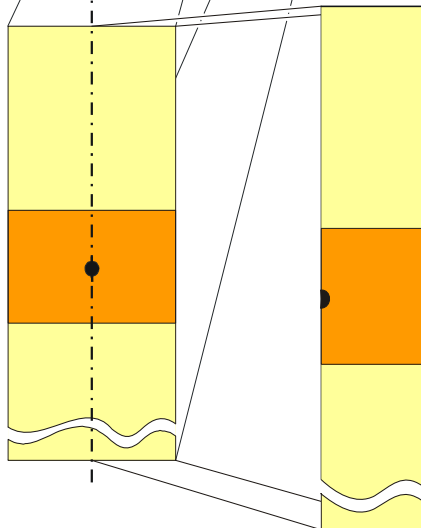
	Clay	Aquifer (sand)
T (thermal)	$\frac{\partial}{\partial t}(\rho_b c_{p,b} T) = \nabla \cdot (\lambda \nabla T) + q$ $\rho_b c_{p,b} = \eta \rho_w c_{p,w} + (1 - \eta) \rho_s c_{p,s}$	$\frac{\partial}{\partial t}(\rho_b c_{p,b} T) + \nabla \cdot (\rho_w c_{p,w} T \mathbf{u}) = \nabla \cdot (\lambda \nabla T)$ <p style="text-align: center;">u = 0</p>
H (hydro)	$\frac{\partial p}{\partial t} = \alpha_H \frac{\partial^2 p}{\partial z^2} + \Lambda \frac{\partial T}{\partial t}$ $\Lambda = \left(\frac{\partial p}{\partial T} \right)_{undrained, oedometer}$	$\frac{\partial}{\partial t}(\eta \rho_w) = \nabla \cdot (\rho_w \mathbf{u})$ <p>with $\mathbf{u} = \frac{k}{\mu} (\nabla p - \rho_w \mathbf{g})$ (Darcy)</p>
M (mech)	$\varepsilon_z = \frac{\Delta p + \beta_d K_d \Delta T}{\lambda_d + 2G}$	$\varepsilon_z = \frac{\beta_d K_d \Delta T}{\lambda_d + 2G}$

Thermal evolution, boundary conditions & mesh

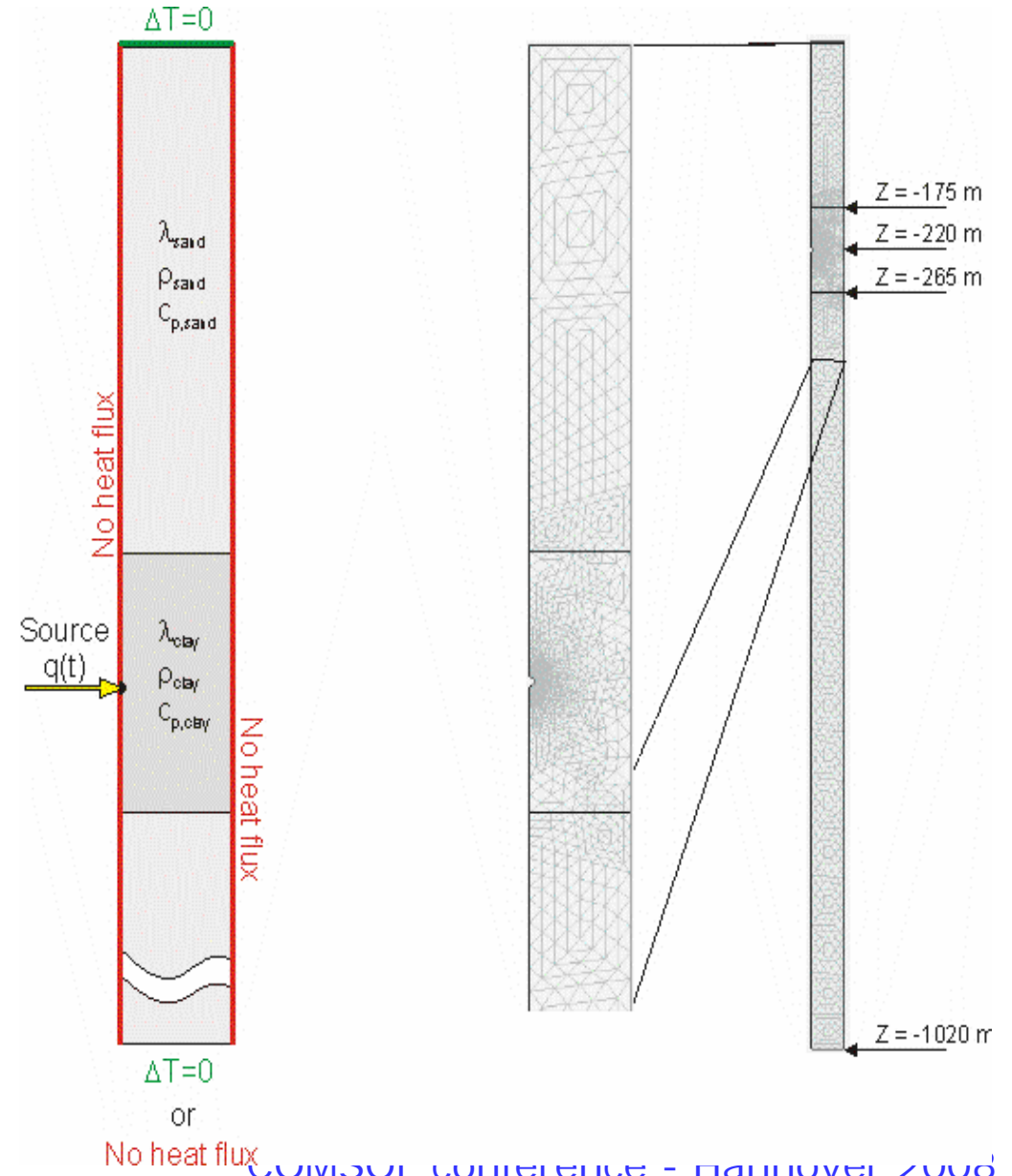
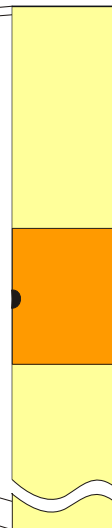
a) 3D world



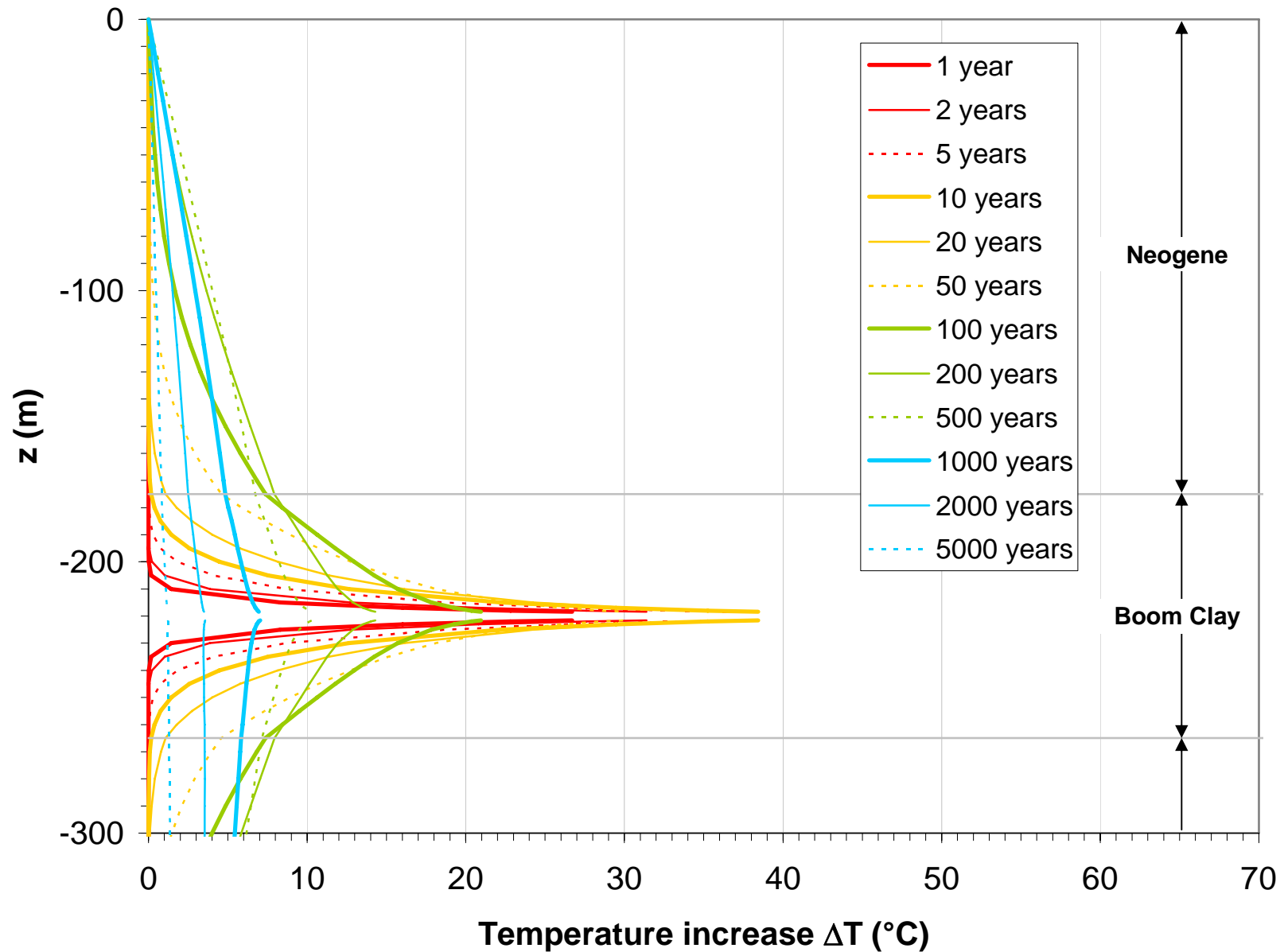
b) 2D model



c) 1/2 2D model



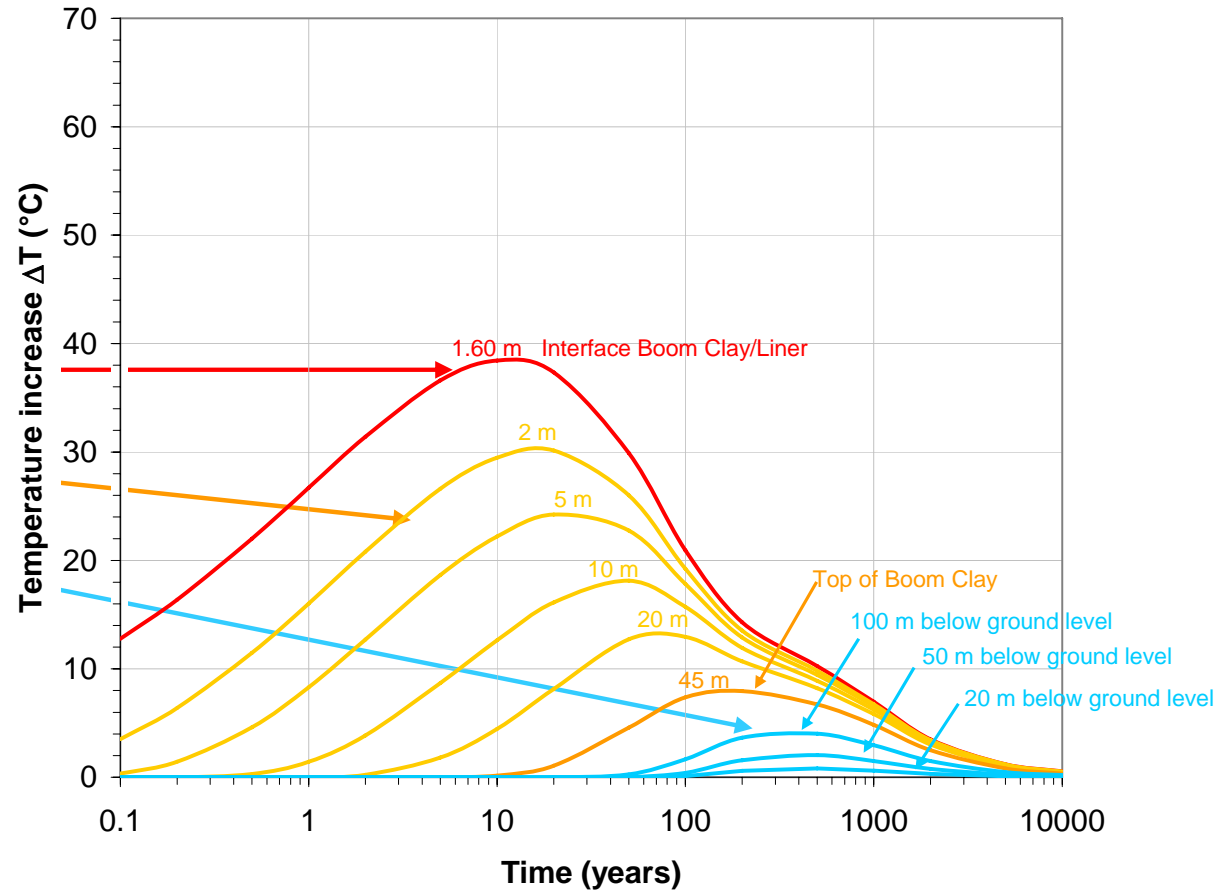
Thermal evolution, Vertical ΔT profiles



Thermal evolution, History of temperature increase ΔT

How **hot** will it be ?

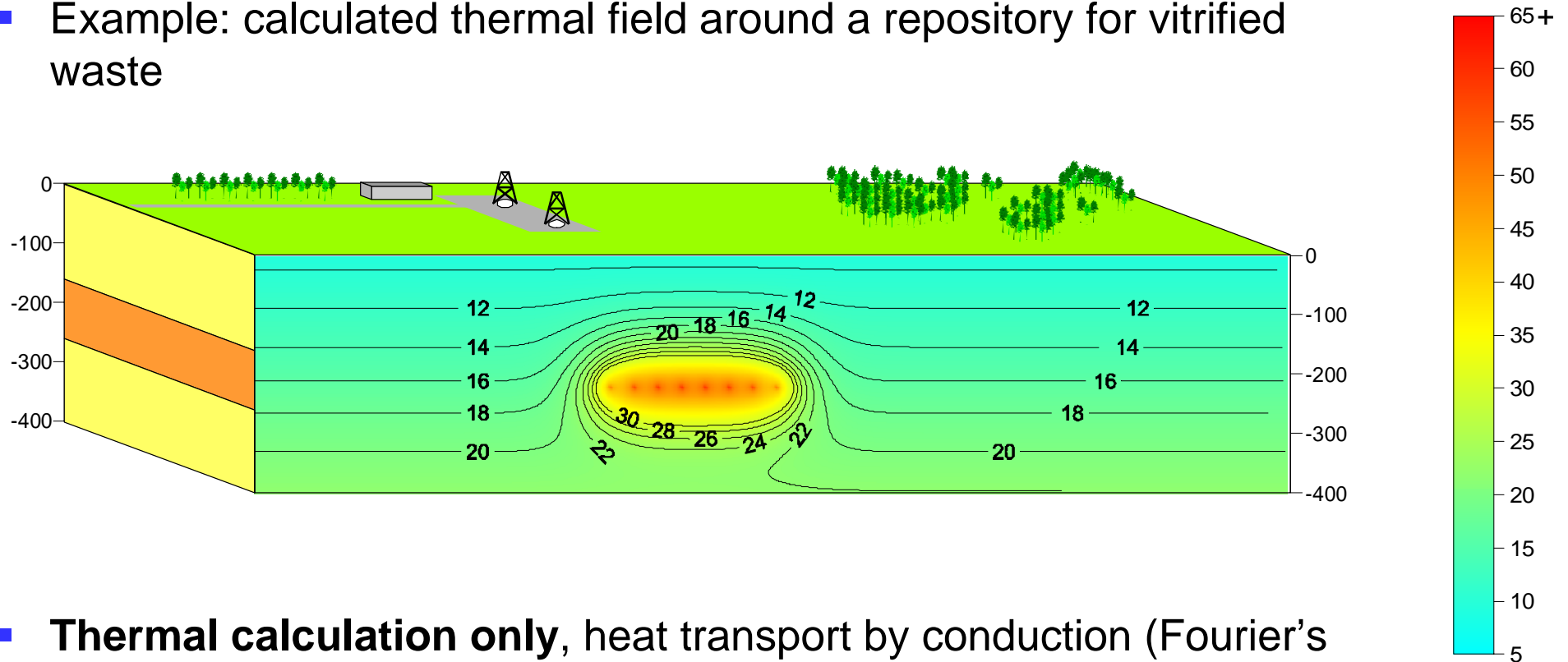
- Waste
- Engineered Barriers System (EBS)
- Clay
- Aquifers



Thermal evolution, full repository Typical results, T contours

How hot will it be ?

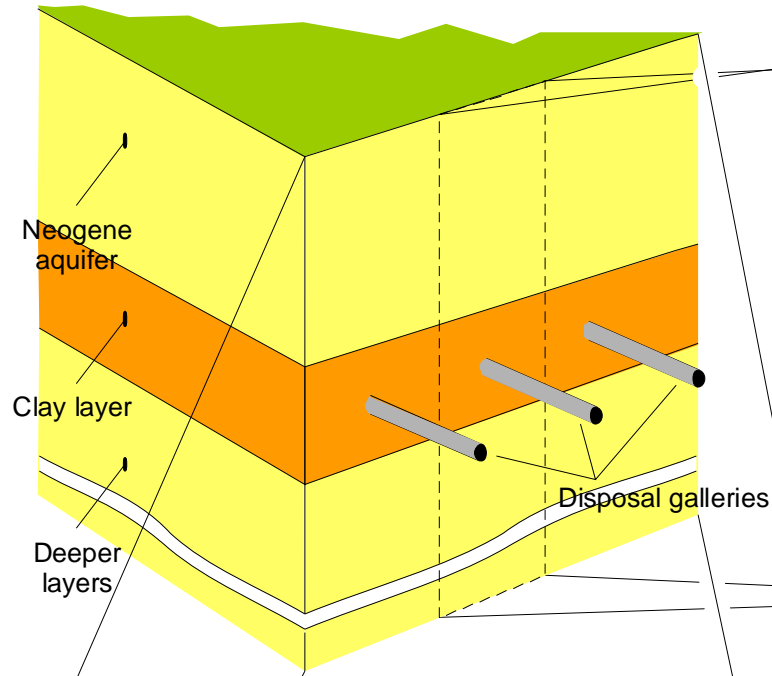
- Example: calculated thermal field around a repository for vitrified waste



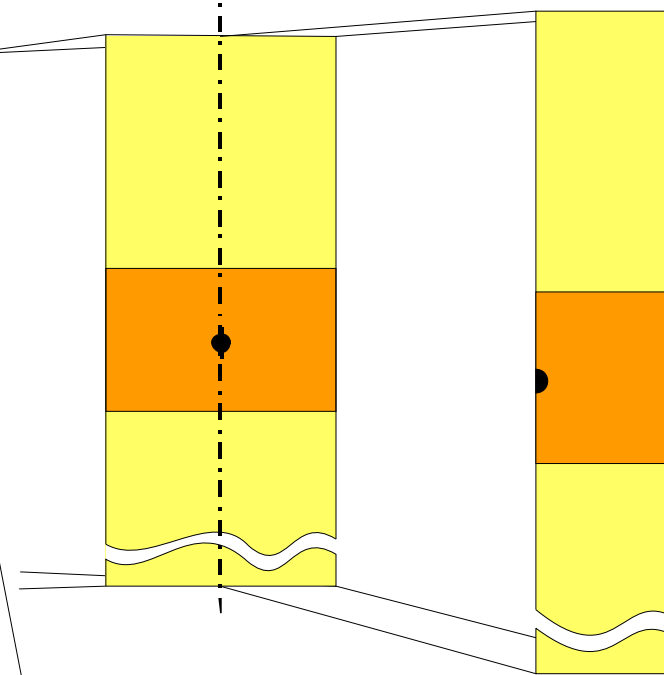
- **Thermal calculation only**, heat transport by conduction (Fourier's law). **Temperature field 100 years after disposal**

Reference geometry T, T-H & T-H-M model reduction

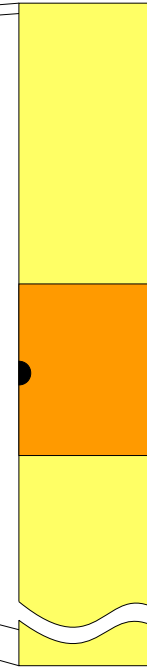
a) 3D world



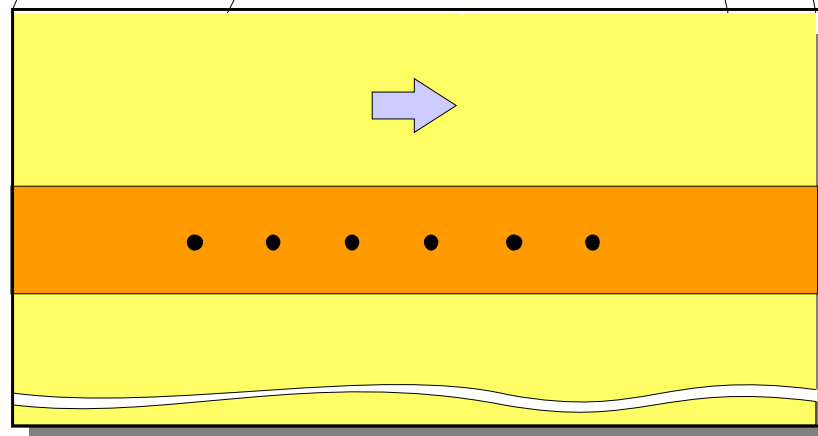
b) 2D model, T only



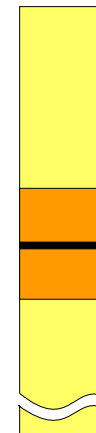
c) 1/2 2D model



d) 2D model, T-H in aquifer



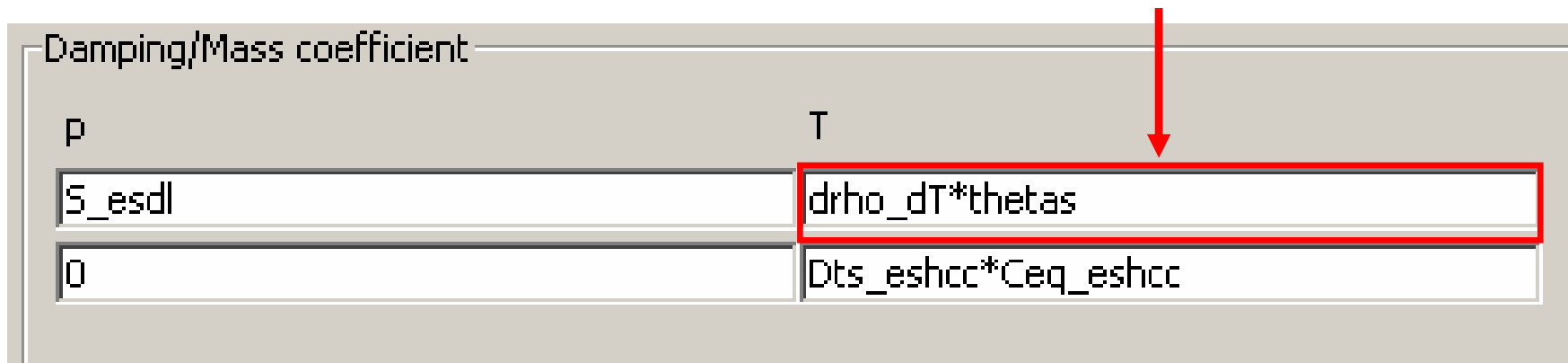
e) 1D T-H-M model



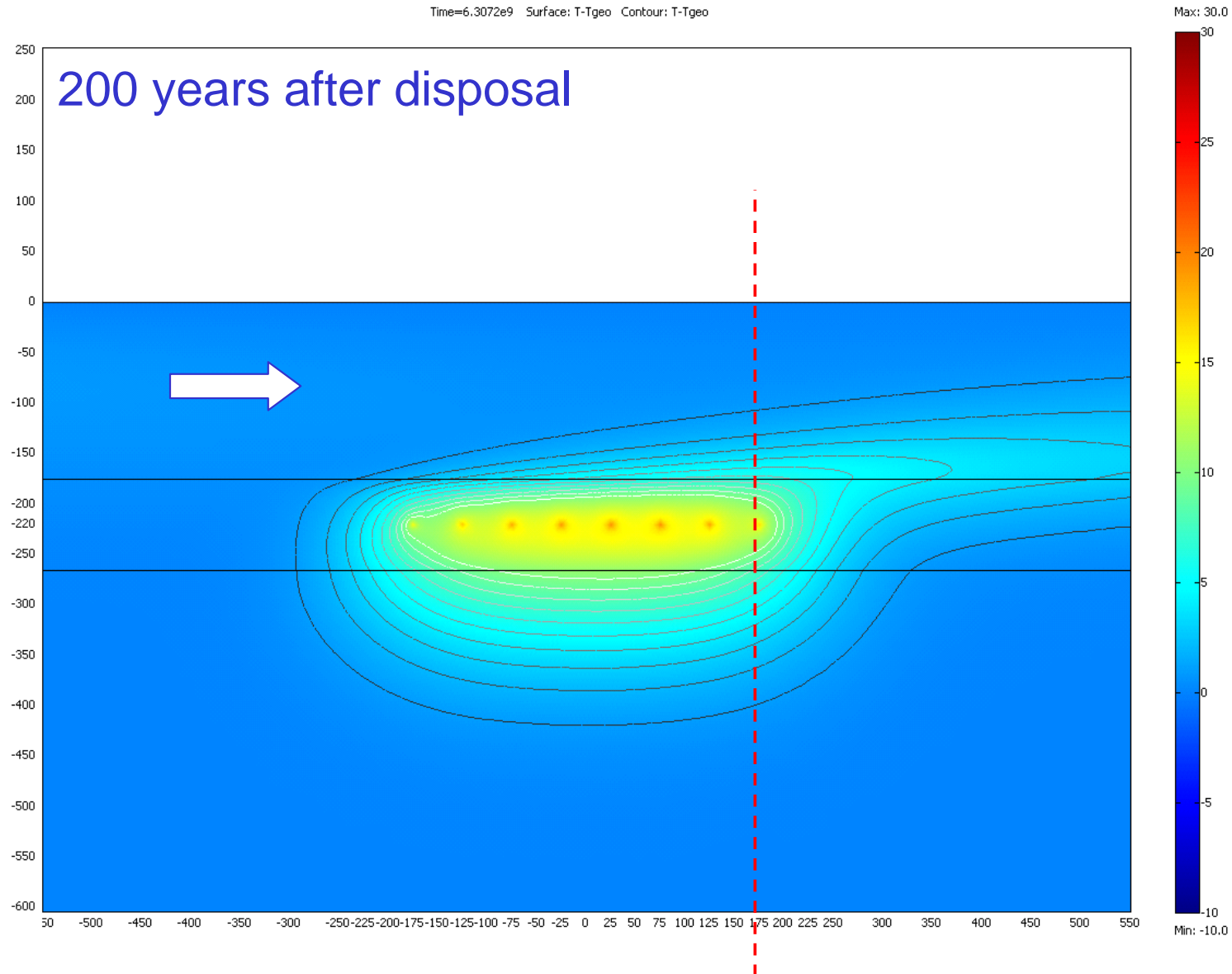
	Clay	Aquifer (sand)
T (thermal)	$\frac{\partial}{\partial t}(\rho_b c_{p,b} T) = \nabla \cdot (\lambda \nabla T) + q$ $\rho_b c_{p,b} = \eta \rho_w c_{p,w} + (1 - \eta) \rho_s c_{p,s}$	$\frac{\partial}{\partial t}(\rho_b c_{p,b} T) + \nabla \cdot (\rho_w c_{p,w} T \mathbf{u}) = \nabla \cdot (\lambda \nabla T)$
H (hydro)	$\frac{\partial p}{\partial t} = \alpha_H \frac{\partial^2 p}{\partial z^2} + \frac{\partial T}{\partial t}$ <p>Very low k, no convective heat transport $\left(\frac{\partial p}{\partial T}\right)_{undrained, oedometer}$</p>	$\frac{\partial}{\partial t}(\eta \rho_w) = \nabla \cdot (\rho_w \mathbf{u})$ <p>with $\mathbf{u} = \frac{k}{\mu} (\nabla p - \rho_w \mathbf{g})$ (Darcy)</p>
M (mech)	$\varepsilon_z = \frac{\Delta p + \beta_d K_d \Delta T}{\lambda_d + 2G}$	$\varepsilon_z = \frac{\beta_d K_d \Delta T}{\lambda_d + 2G}$

COMSOL Multiphysics implementation and auxiliary equations

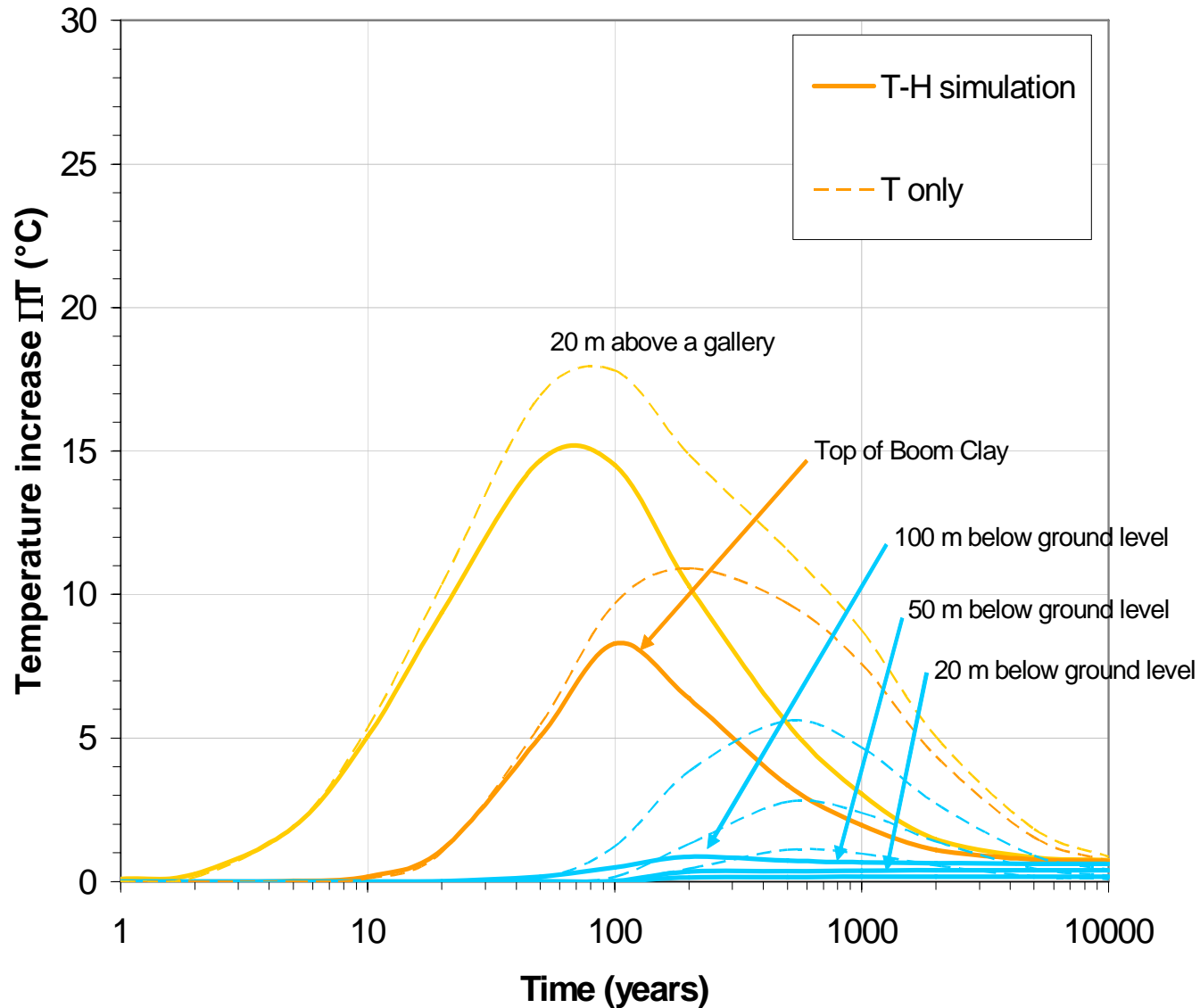
- Use of Earth Science Module (convenient, but not required)
 - **H**: Darcy's law (**esdl**)
 - **T**: Conduction & convection in porous media (**eshcc**)
 - Water density: $\rho = 1000.2 - 0.005 \times T^2$ [kg/m³] (T in °C)
 - Water viscosity: $\mu = \rho \cdot 9.2 \times 10^{-7} \cdot \exp(2050 / (273.15 + T))$ [Pa·s]
- No convection in low-permeability clay & geological layers below
 - Simply do not solve for flow in these subdomains ☺
- Coupling of heat and flow equations:
 - **H**→**T**: Use velocities from **esdl** in **eshcc**
 - **T**→**H**: COMSOL > Physics > Equation system > Subdomain settings



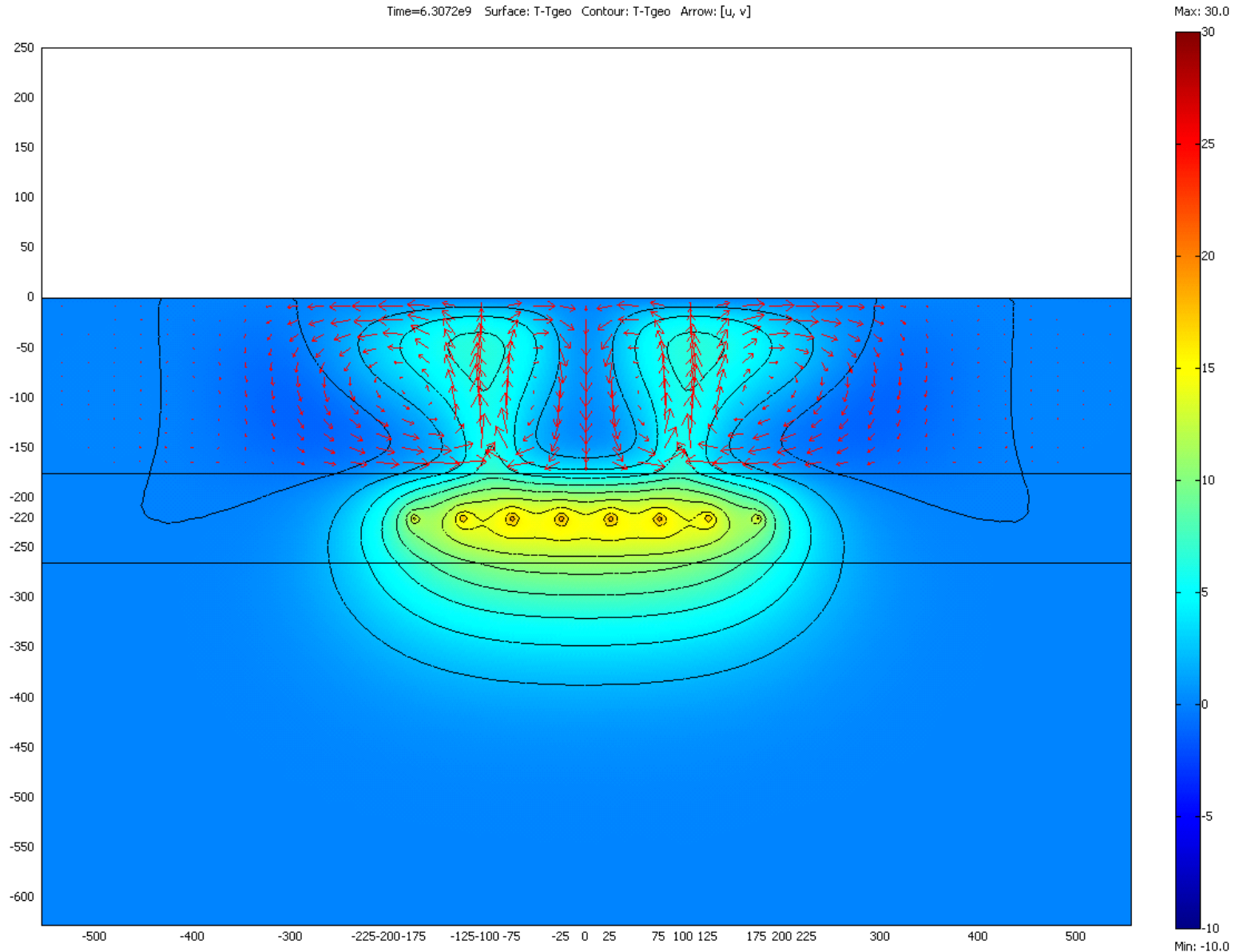
T-H evolution, effect of local flow pattern



T-H evolution, effect of local flow pattern



T-H evolution, convection cells only in the absence of base flow !



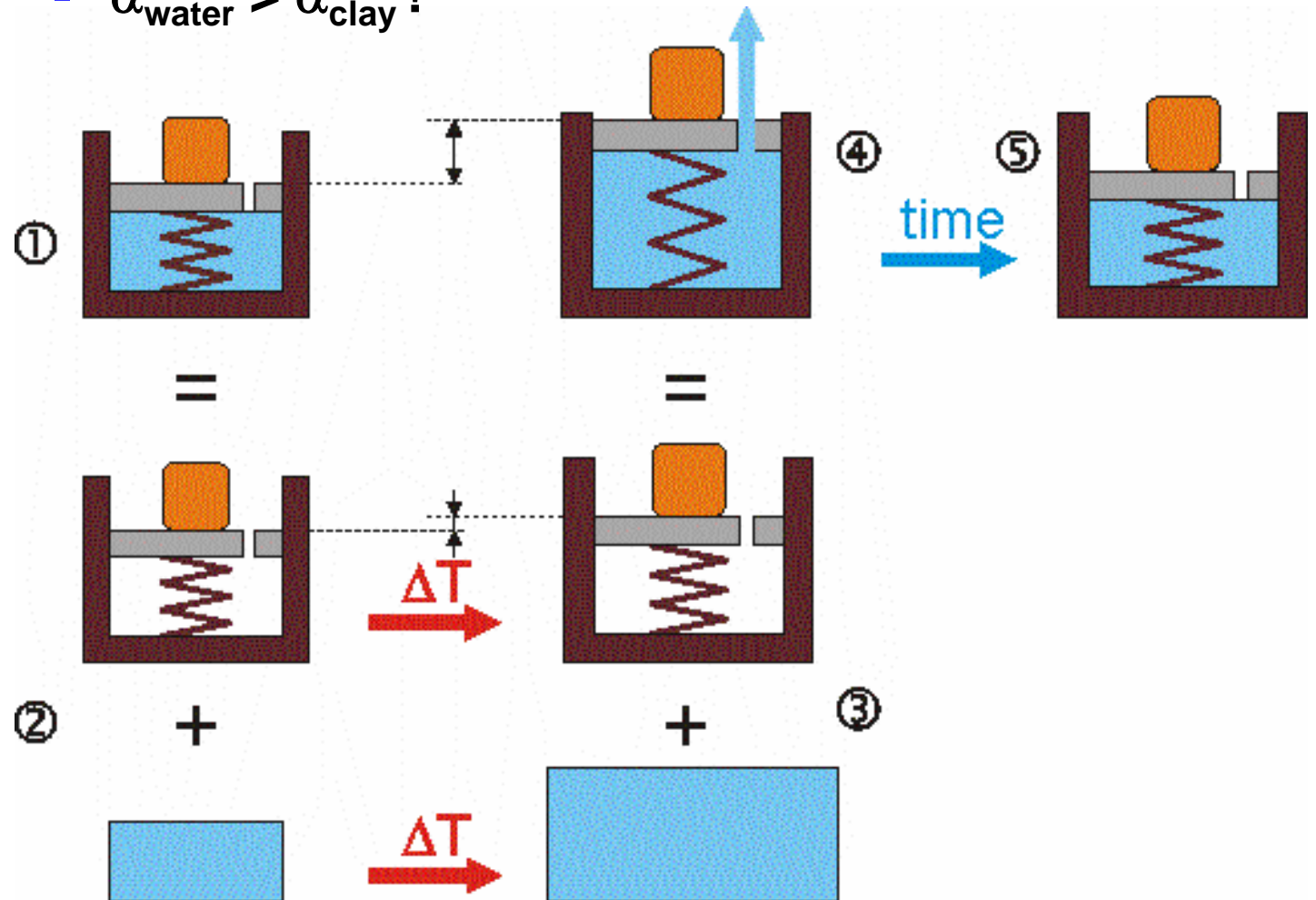
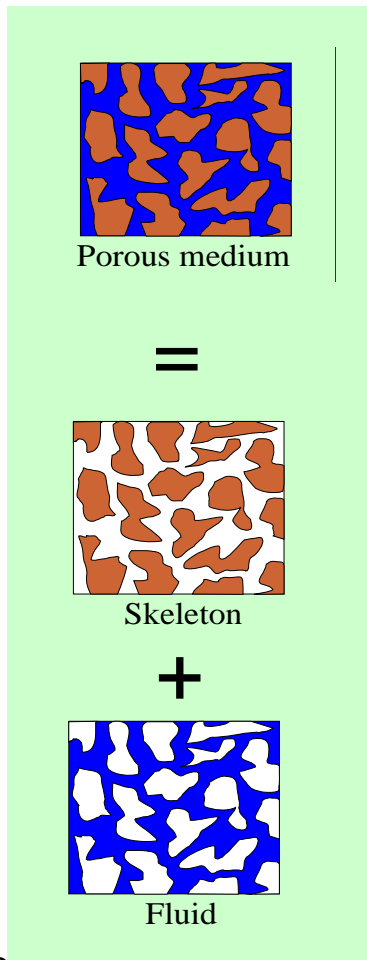
- Cause of uplift: **thermal expansion**

Material	Expansion coeff. ($\text{m}^3/\text{m}^3 \text{ } ^\circ\text{C}^{-1}$)	Symbol
Clay, drained	3×10^{-5}	β_d
Clay, undrained	13×10^{-5}	β_u
Water	21×10^{-5}	β_w
Sand, drained	3×10^{-5}	β_d

- **Aquifers**: excess water volume can quickly be accommodated
- **Clay**: overpressures, which slowly dissipate

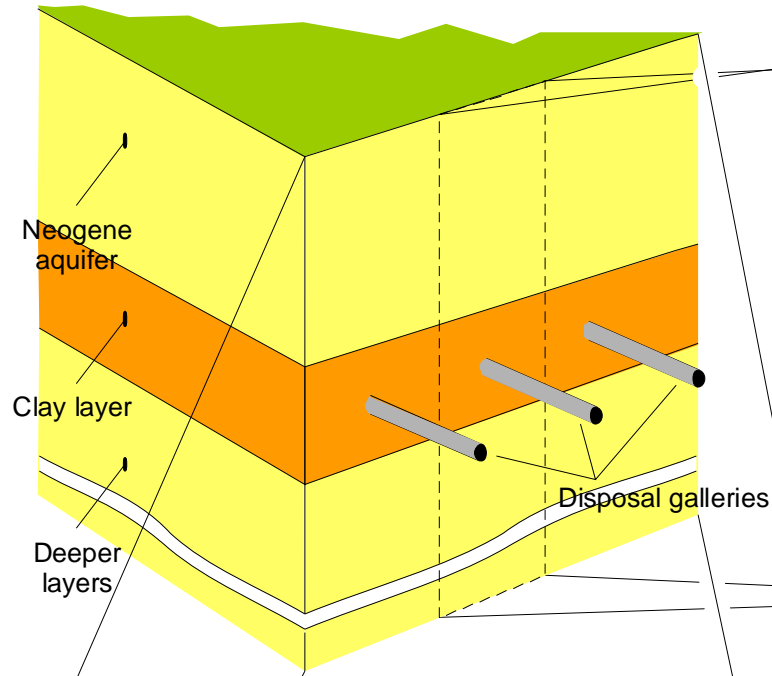
Terzaghi's analogy adapted to T → HM

- $\Delta T \rightarrow$ thermal expansion (α)
- $\alpha_{\text{water}} > \alpha_{\text{clay}}$!

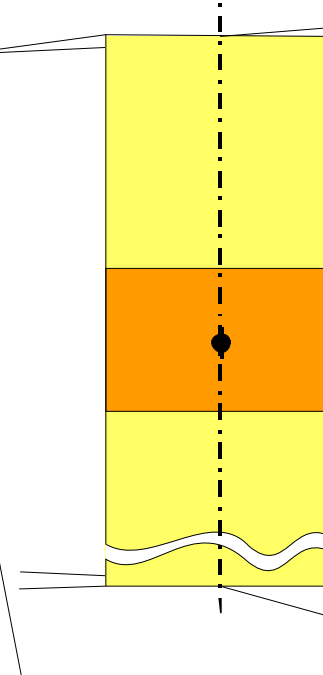


Reference geometry T, T-H & T-H-M model reduction

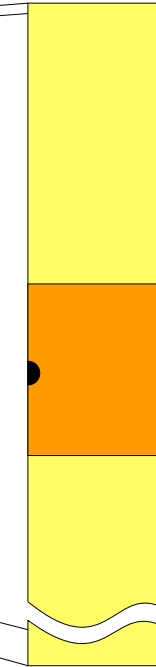
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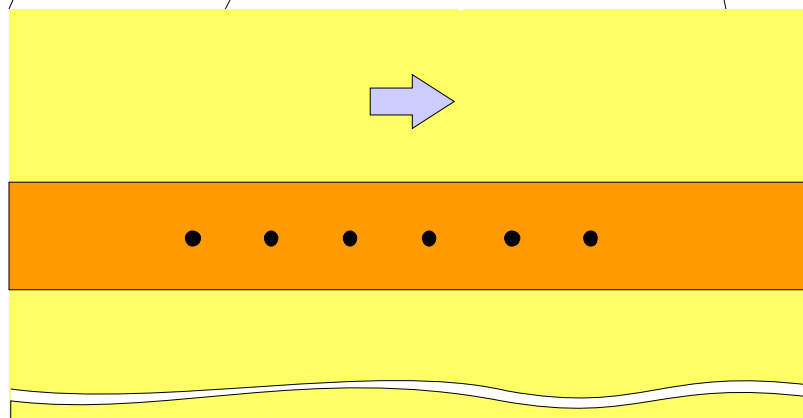
b) 2D model, T only



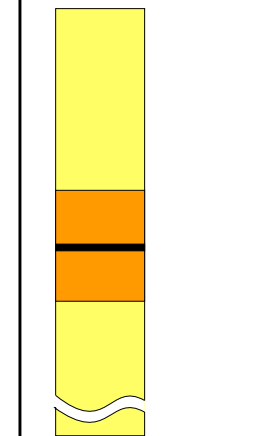
c) 1/2 2D model



d) 2D model, T-H in aquifer



e) 1D T-H-M model



Model equations: **T-H-M**

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H (hydro)	$\frac{\partial p}{\partial t} = \alpha_H \frac{\partial^2 p}{\partial z^2} + \Lambda \frac{\partial T}{\partial t}$ $\Lambda = \left(\frac{\partial p}{\partial T} \right)_{undrained, oedometer}$	$\frac{\partial}{\partial t}(\eta \rho_w) = \nabla \cdot (\rho_w \mathbf{u})$ <p style="text-align: center;">1D, heat transport by conduction</p> <p>with $\mathbf{u} = \frac{k}{\mu} (\nabla p - \rho_w \mathbf{g})$ (Darcy)</p>
M (mech)	$\varepsilon_z = \frac{\Delta p + \beta_d K_d \Delta T}{\lambda_d + 2G}$	$\varepsilon_z = \frac{\beta_d K_d \Delta T}{\lambda_d + 2G}$

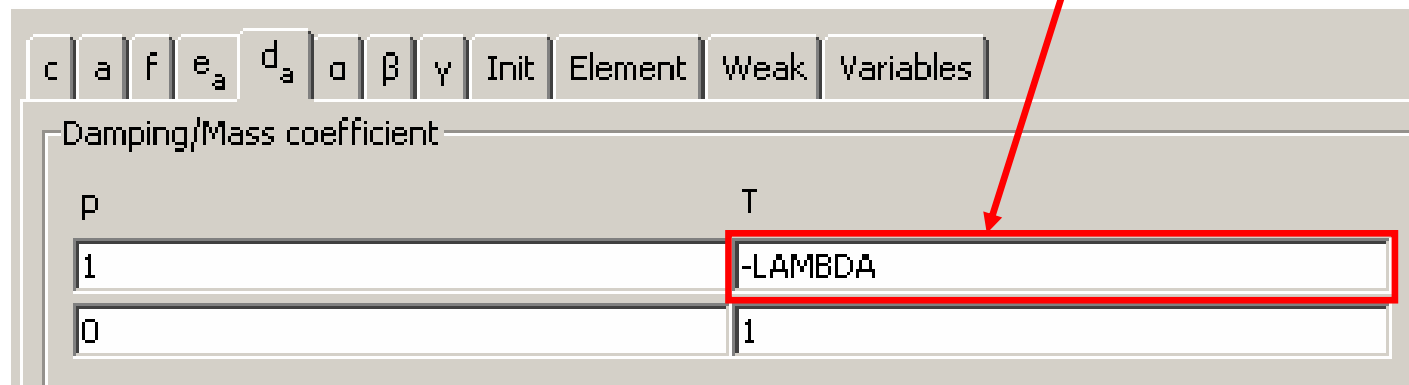
- Summary of model equations (details in Picard & Giraud, 1995)

- Heat transport:
$$\frac{\partial T}{\partial t} = \alpha_T \frac{\partial^2 T}{\partial z^2} + \frac{q}{\rho_b c_{p,b}}$$

- Porewater pressure dissipation:
$$\frac{\partial p}{\partial t} = \alpha_H \frac{\partial^2 p}{\partial z^2} + \Lambda \frac{\partial T}{\partial t}$$

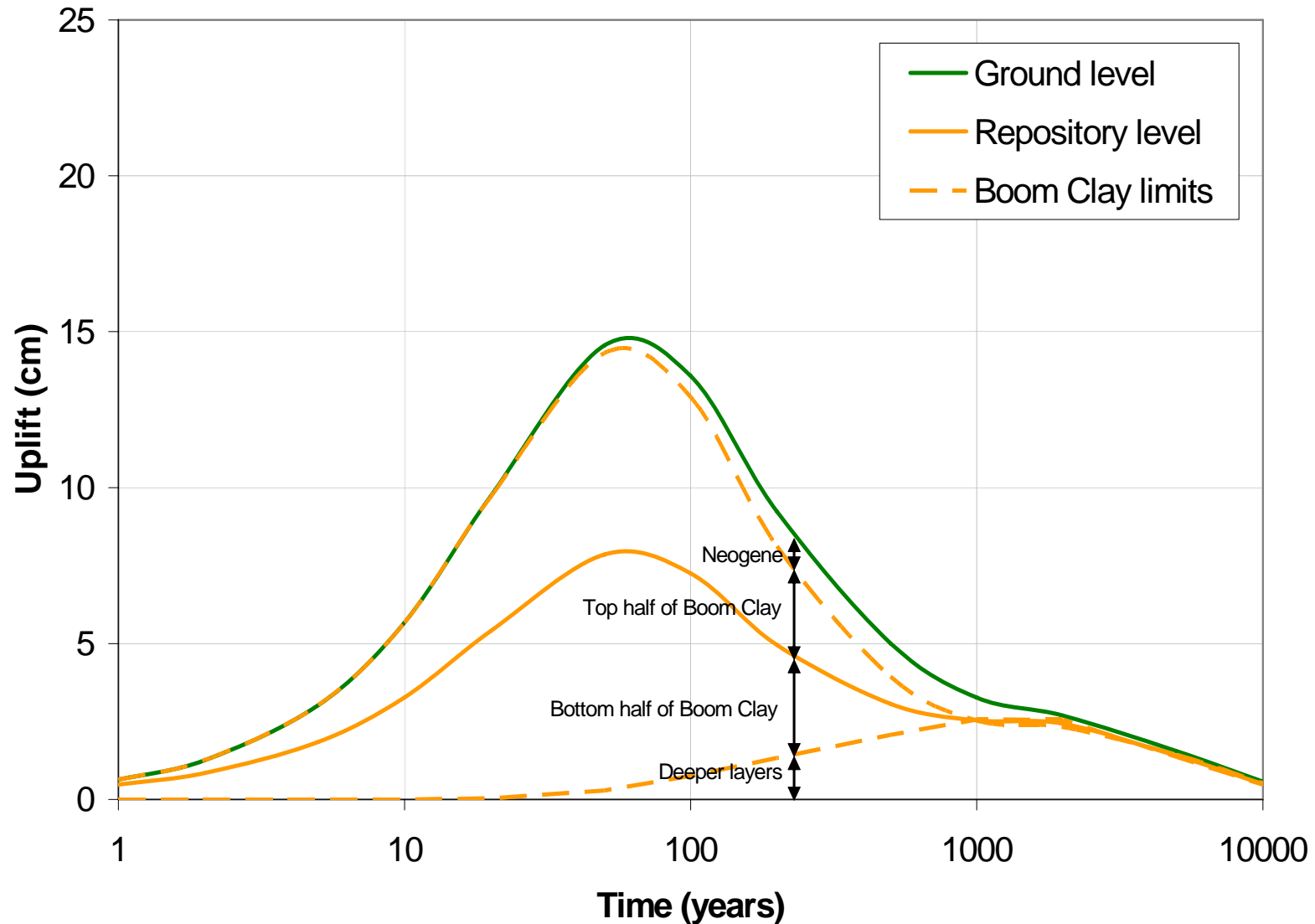
- Vertical deformation:
$$\varepsilon_z = \frac{\Delta p + \beta_d K_d \Delta T}{\lambda_d + 2G}$$

- Solve two 1D diffusion equations, then integrate ε_z over depth
- "Coupling" in COMSOL Multiphysics:
 - COMSOL > Physics > Equation system > Subdomain settings



Uplift evolution

note that most of the uplift is due to thermal expansion of poorly drained clay (water)



- **Modelling the geological disposal of radwaste**
 - Large time scales
 - Multiple spatial scales (near field, far field)
 - Many processes involved, some of these are strongly coupled
- **Complexity ?**
 - Multidisciplinary rather than intrinsically complex
 - Large uncertainties, emphasize robust modelling (simplifications)
- **How COMSOL Multiphysics fits in the picture**
 - **VERSATILITY**: 1 toolbox, many possible uses in R&D programme
 - **Thermal evolution of the far field** (this presentation)
 - **Phenomenological analysis**: **near field THM**, buffer THMC, **chemo-osmosis**, reactive transport, **unsaturated flow**, multiphase flow,...
 - **Performance Assessment**: **radionuclides release & transport**

Thank YOU for your attention.

Thanks go also to



ONDRAF/NIRAS,

the Belgian National Radioactive Waste Agency,
for continued support & funding.